- Doyle, J. A., Sauquet, H., Scharaschkin, T. and Le Thomas, A., Phylogeny, molecular and fossil dating, and biogeographic history of Annonaceae and Myristicaceae (Magnoliales). *Int. J. Plant Sci.* (*Suppl.*), 2004, **165**, 555–567.
- Raven, P. H. and Axelrod, D. I., Angiosperm biogeography and past continental movements. *Ann. Miss. Bot. Gard.*, 1974, 61, 539-673.
- 34. Morley, R. J., Origin and Evolution of Tropical Rainforests, Wiley, Chichester, UK, 2000.
- Morley, R. J., Interplate dispersal routes for megathermal angiosperms. *Perspect. Plant Ecol. Evol. Syst.*, 2003, 6, 5–20.
- Tiffney, A., Perspectives on the origin of the floristic similarity between eastern Asia and eastern North America. J. Arnold Arboretum., 1985, 66A, 73–94.
- 37. Tiffney, B., The Eocene North Atlantic land bridge: its importance in Tertiary and modern phytogeography of the northern hemisphere. *J. Arnold Arboretum*, 1985, **66B**, 243–273.
- 38. Manchester, S. R., Biogeographical relationships of North American tertiary floras. *Ann. Miss. Bot. Gard.*, 1999, **86**, 472–522.

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Dwarfism and Lilliput effect: a study on the *Glossopteris* from the late Permian and early Triassic of India

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The 'Lilliput effect' represents the phenomenon whereby there is a pronounced reduction in the size of biota associated with the aftermath of mass extinction. This fact has been supported by the evidence of dwarfism both in invertebrates and vertebrates recorded after the end-Permian mass extinction event. The extinct genus *Glossopteris* belonging to seed ferns Glossopteridales is one of the best known fossil taxon that flourished during the Permian and continued its existence till Triassic. In contrast to the Permian, the Triassic was a time when greenhouse conditions with an increased temperature and widespread aridity prevailed as evidenced by the global dataset. The new set of environmental conditions in the Triassic posed a major challenge for the existing *Glossopteris* lineage, whereby the smaller forms (dwarfs) with reduced leaf surface area continued and sustained. The present study from different late Permian and early Triassic formations of India is aimed at unravelling the changes in morphological traits of seven species of *Glossopteris* leaves whose existence continued surpassing the Permian–Triassic mass extinction event.

Keywords: Dwarfism, extinction event, *Glossopteris*, Lilliput effect, Permian–Triassic boundary.

GLOSSOPTERIS is one of the best known taxon belonging to the extinct order of the seed ferns Glossopteridales (family Glossopteridaceae). It is one of the first fossil plants described and named by Brongiart^{1,2}. The name implies lanceolate to tongue-shaped entire leaves, which are characterized by a prominent midrib and reticulate secondary venations^{2–5}. It has been extensively recorded from the once united southern hemisphere called 'Gondwana', which comprised of India, Australia, New Zealand, South America, Africa and Antarctica.

The origin of the glossopterids took place in the southern hemisphere around the beginning of the Permian ~ 290 Ma (ref. 6). *Glossopteris* is considered as the stratigraphic marker for the Permian throughout all the Gondwana continents. Its existence continued up to the Triassic and there is one report of its doubtful occurrence in the Jurassic⁷.

The late Permian megafossil assemblage of India (Bijori Formation, Kamthi Formation and Raniganj Formation) is dominated by larger-sized species of *Glossopteris*^{8–15}. They were broadleaved and presumed to be deciduous, capable of living in a wide range of fluvio-lacustrine subenvironments^{12,13,16,17}. In contrast, species of *Glossopteris* recorded from the early Triassic (Panchet Formation) were comparatively smaller in dimension than the underlying Permian¹⁸. The leaves of early Triassic sediments had a reduced surface area which was adapted to withstand the extreme climatic condition^{19–21}.

The 'Lilliput effect' represents the phenomenon whereby there is a pronounced reduction in the size of the biota associated with the aftermath of mass extinction event^{22–25}. After a particular extinction event, the biotic crisis is normally represented in two major ways: (i) the larger forms which are structurally specialized species tend towards extinction^{26,27} and (ii) species whose size decreases (dwarfism) tend to continue^{28,29}. The 'Lilliput effect' has been established for the end-Permian Mass Extinction Event (EPME). This fact is supported by the evidence of dwarfism of invertebrates^{30,31} and relatively few records of vertebrates^{32–34}. A single factor cannot be held responsible for the phenomenon of dwarfism, i.e. 'Lilliput effect'. Factors that are responsible for this phenomenon can be either biotic or abiotic or a combination of

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both (Figure 1) which act in a cumulative way^{29,35-37}. Abiotic factors that are responsible for the same are: mean temperature rise, anoxia, increased CO₂ concentration in the atmosphere, edaphic factors, climatic change, loss of provinciality and ecological niche. Biotic factors that are the causative agent for the phenomenon are symbiotic dissociation, diseases, reduction in primary productivity, nutrient scarcity, reduced vigour and intra- or intergeneric and specific competition. However, the most important factor for dwarfism, at the Permian-Triassic Boundary can be attributed to increased aridity, higher CO₂ concentration (about 2000 ppm) in the atmosphere and substantial rise in temperature $(\sim 8^{\circ}C)^{38}$. A similar reduction in leaf width has been observed on wellpreserved fossil leaves from the terrestrial Triassic-Jurassic sections in East Greenland and Scania, Sweden³⁹. Comprehensive data generated on Lepidopteris-Thaumatopteris flora (Rhaetian-Hettangian) reveal the extent of high-temperature injury to leaves across the Triassic-Jurassic Boundary. The authors opine that volcanicity associated with the Central Atlantic Magmatic 199 ± 2.4 Ma) Province (CAMP, induced 'supergreenhouse' conditions resulting in a fourfold increase in atmospheric CO₂ and 3-4°C warming across the Triassic–Jurassic Boundary. The authors³⁹ further suggest that reduction in leaf size was to avoid lethally high leaf temperature. Increased SO₂ levels resulting from massive volcanism in the CAMP have also been implicated as a factor in reduced leaf size⁴⁰. Experiments conducted on the nearest living equivalent taxa of the Triassic-Jurassic Greenland flora have revealed that fumigation by SO₂ resulted in leaf roundness as observed in the fossil flora. Therefore, rising SO₂ levels as a consequence of volcanism during the Permian-Triassic extinction event (PTEE) could also be a factor in reduced leaf size.



Figure 1. Possible biotic and abiotic factors responsible for 'Dwarfism and Lilliput effect'.

The PTEE that occurred ~ 252 Ma is the most severe and lethal event in the Earth's history. It witnessed the extinction of about 90% of all marine species and 70% of terrestrial species. The magnitude of floral extinction along with the overall change in diversity is more ambiguous. There are no major peaks in mass extinction for plants in the fossil record, excepting the Permian-Triassic Boundary⁴¹. Studies have revealed that the ecological instability in plant communities and their ultimate extinction was before or coincident with the peak of faunal extinction^{31,42,43}. But, the effect on plants was much lesser compared to the marine and faunal records. However, some plants became extinct at the Permian-Triassic Boundary. Evidence shows that at the Permian-Triassic Boundary there was a sudden increase in pumping of massive volume of CO₂ into the atmosphere that led to the global warming and short-term production of acid rain⁴⁴. This major catastrophic event affected the existing plant life. However, Glossopteris is the only plant taxon that could survive the event and is found beyond the Permian-Triassic Boundary in relative abundance. In fact, Glossopteris is an exceptional form that overcame the Permian-Triassic boundary mass extinction event and continued with reduced vigour in the early Triassic. Other associated late Permian plant taxa, viz. Vertebraria, Gangomopteris, Phyllotheca, Dichotomopteris, Neomariopteris, etc. could not sustain the extreme climatic conditions during the Permian-Triassic Boundary crisis.

In the present study, an attempt has been made to decipher the changes in the various aspects of morphological traits of the species of *Glossopteris* from the late Permian and early Triassic succession of India. The abovementioned formations with their respective geographical locations are shown in Figure 2.

The late Permian species of Glossopteris were studied from three different formations namely Bijori, Kamthi and Raniganj^{8–15}; while for early Triassic, the Panchet Formation was taken into consideration^{19–21}. The late Permian-Triassic sequences in different Gondwana basins of Peninsular India have been listed in Table 1. Lithology of the late Permian sediments of Raniganj Formation is characterized by predominantly fine-grained, occasionally feldspathic, micaceous, sometimes argillaceous, greenish-grey to whitish-grey sandstone. The sandstones of Raniganj Formation are frequently interbedded with micaceous siltstone. Grey as well as carbonaceous shales and coal seams are well developed. Lithologically the Bijori Formation comprises sandstones with micaceous shale and carbonaceous shales along with coal streaks. The lithological features as well as floral composition establish its homotaxis with the Raniganj Formation. The Kamthi Formation is predominantly composed of conglomerates, usually soft and argillaceous grits, finegrained sandstone and shales. The floral assemblage is similar to that of the Raniganj Formation. The rocks of the early Triassic sediments of Panchet Formation are



Figure 2. Map showing the different formations of Gondwana Basin^{67,68} in Peninsular India from where the present analyses have been carried out.

| Stand | dard | Gondwana basins | Damodar Valley Basin | Koel Valley Basin | Rajmahal Basin | South Rewa Basin | Satpura Basin | Pranhita– Godavari Valley Basin | Mahanadi Valley Basin |
|-------|---------|--------------------|----------------------------|-------------------------|------------------------|----------------------|---------------------------------|---------------------------------------|-----------------------------|
| Lowe | er Jura | ssic | | | | Hartala Hill Beds | | Dharmaram | Pathargarh |
| | r | Rheatian | | | | Tiki Formation | | Formation | Beds |
| | Uppe | Norian | | | Dubrajpur Formation | | | Maleri Formation | |
| ssic | | Carnian | Supra-Panchet | Mahadeva | | Parsora Formation | Bagra Beds | Bhimaram | |
| Tria | ldle | Ladinian | Formation | Formation | | | Denwa | Formation | |
| | Mid | Anisian | | | | Nidpur Beds | Formation Pachmari | Yerrapalli Formation | |
| | Lower | Sythian | Panchet Formation | Panchet Formation | | Panchet Formation | Formation Almod Formation | Mangli Beds | Sarimunda Hill Beds |
| Late | Permia | an | Raniganj Formation | Raniganj Formation | | Pali Formation | Bijori Formation | Kamthi Formation | Kamthi Formation |

Table 1. The late Permian-Triassic sequences in different Gondwana basins of Peninsular India

 Table 2.
 Number of specimens of Glossopteris species analysed

| - | - | |
|----------------------------|--|--|
| Species of Glossopteris | Number of late Permian species studied | Number of early Triassic species studied |
| Glossopteris angustifolia | 8 | 6 |
| Glossopteris browniana | 10 | 4 |
| Glossopteris communis | 12 | 8 |
| Glossopteris conspicua | 14 | 4 |
| Glossopteris damudica | 12 | 2 |
| Glossopteris indica | 10 | 6 |
| Glossopteris retifera | 8 | 4 |
| | | |

Total number of specimens studied (n) = 108. Seventy four specimens from the late Permian. Thirty four specimens from the early Triassic.

CURRENT SCIENCE, VOL. 107, NO. 10, 25 NOVEMBER 2014

clearly distinguishable from the underlying Raniganj Formation by the characteristic red shales, rich micaceous sandstones and absence of coal and carbonaceous shales. The lowermost beds of the Panchet Formation comprise shales and thin beds of greenish-brown mudstones, interbedded with micaceous greenish-grey siltstones⁴⁵.

Materials for the present study consisted of the illustrated, type and figured specimens of the *Glossopteris* species housed in the repository of Birbal Sahni Institute of Palaeobotany (BSIP), Lucknow, India. In all, 108 specimens preserved as impressions have been taken into consideration for the present analysis (Table 2). Based on various morphological traits of the leaf architecture,

| | - | ; | | | | | ۰ ۲ | | | | | | |
|------------------------------|---|-------------------------------------|------------------------------------|---|--------|-----------------------------|--------|---|------------------|--|--|-------------------------------|-------------------------------|
| | | Siz | e | I | | | | Midril | ~ | | Ave | erage concenti | ration of veins |
| Species of Glossopteris | Age/ formations | Maximum available length (cm) | Maximum available width (cm) | Shape | Apex | Base | Margin | Type | hickness (mm) | Angle of emergence of secondary veins | Mesh type | Towards midrib (per cm) | Towards margin (per cm) |
| Glossopteris angustifolia | Late permian (Bijori/ Kamthi/ Raniganj formations) | 13.0 | 2.00 | Lorate- oblong to narrow- oblanceolate | Obtuse | Acute- cuneate | Entire | Medium to broad, flat, striated, evanescent towards apex | 5.00 | Emerge at 10–15°, fork once or more from midrib to margin | Elongated, oblong to linear- polygonal | 18–25 | 12–15 |
| | Early Triassic (Panchet Formation) | 8.50 | 2.00 | Oblanceolate to spathulate | Acute | Tapering towards base | Entire | Distinct, finely striated | 2.00 | Emerge at 15–20°, near base parallel to midrib | Broad, oblong, narrower near margin | 20–30 | 15–23 |
| Glossopteris browniana | Late Permian (Bijori/ Kanthi/ Rantagnj formations) | 16.4 | 5.60 | Narrow oblong | Acute | Acute- cuneate | Entire | Thick at base, tapers gradually, distinct at apex | 2.50 | Emerge at 20-45°, slightly arched, reach the margin at an open ar an open ar an open ar an open ar an open arched. | Narrow- broad, elongated polygonal meshes | 22-40 | 25-45 |
| | Early Triassic (Panchet Formation) | 7.40 | 5.00 | Lanceolate to spathulate | Acute | Tapering towards base | Entire | Distinct, flat, longitudinally ribbed, persist up to apex | 1.50 s | angre Emerge at 35-45°, dichotomizing and anastomosing | Oblong- polygonal, narrower towards margin | 10–12 | 18–20 |
| Glossopteris communis | Late Permian (Bijori/ Kamthi/ Raniganj formations) | 25.0 | 6.00 | Narrow elliptic to oblong lanceolate | Acute | Acute- cuneate | Entire | Thick, flat, persists up to apex | 3.00 | Emerge at 10–18°, arching backwards | Long, narrow, oblong, polygonal | 32-40 | 45-60 |

1738

CURRENT SCIENCE, VOL. 107, NO. 10, 25 NOVEMBER 2014

| Size Maximum available 11.0 11.0 | | | | Mor | rphologica | I characters of | leaf | | | | |
|--|------------------------------------|-----------------------------|--------------------|-----------------------------|------------|--|----------------|--|---|-------------------------------|-------------------------------|
| Maximum available length (cm) 11.0 | | | | | | | ÷ | | Aver | rage concentr | ation of veins |
| 11.0 | Maximum available width (cm) | Shape | Apex | Base | Margin | Type | Thickness (mm) | Angle of emergence of secondary veins | Mesh type | Towards midrib (per cm) | Towards margin (per cm) |
| | 4.60 | Lanceolate to spathulate | Acute | Tapering towards base | Entire | Distinct, flat. tapering towards apex | 2.50 | Emerge at 30–35°, dichotomizing and anastomosing | Narrow- elongated, equal width throughout the lamina | 24-26 | 34–36 |
| 11.0 | 4.20 | Oblong to oblanceolate | Indis- cernible | Indis- cernible | Entire | Flat, persistent | 2.00 | Emerge at 20–35°, straight course to margin | Long, broad, pentagonal to hexagonal, uniform | 7-10 | 9–15 |
| 6.50 | 3.60 | Obovate | Acute | Tapering towards base | Entire | Distinct, evanescent towards apex | 1.50 | Emerge at 35–40°, travel straight to margin | Large, broad, oblong- polygonal, narrower near margin | 4–6 | 6-8 |
| 18.2 | 6.30 | Elliptic | Obtuse | Not preserved | Entire | Thick, persistent, evanescent towards apex | 5.00 | Emerge at 50–70°, straight course to margin | Broad, polygonal, narrow near margin, short near midrib | 10–13 | 30–32 |
| 3.00 | 2.70 | Indiscernible | Indis- cernible | Indiscer- nible | Entire | Distinct, strong | 1.25 | Emerge at 45°, after emergence bends sharply and continues up to margin | Oblong- polygonal | 4-6 | 9-14 |

CURRENT SCIENCE, VOL. 107, NO. 10, 25 NOVEMBER 2014

RESEARCH COMMUNICATIONS

| Table 3. (Lo | (| | | | | Mo | rnhologica | al characters of] | eaf | | | | |
|----------------------------|---|--------------------------|-------------------------|-------------------------------------|-----------------------|----------------------------------|------------|---|-------------------|--|--|--------------------|--------------------|
| | | Siz | ze | | | | | | | | Ave | srage concenti | ation of veins |
| | | Maximum | Maximum | | | | | Midri | Q | <u> </u> | | Towards | Towards |
| Species of Glossopteris | Age/ formations | available length (cm) | available width (cm) | Shape | Apex | Base | Margin | Type | Thickness (mm) | emergence of secondary veins | Mesh type | midrib (per cm) | margin (per cm) |
| Glossopteris indica | Late Permian (Bijori/ Kamthi/ Raniganj formations) | 31.0 | 3.50 | Narrow elliptic | Acute | Acute | Entire | Distinct, persists up to apex | 5.00 | Emerge at 12–20°, veins anastomosing | Broad near midrib, narrow towards margin | 22-35 | 35–48 |
| | Early Triassic (Panchet Formation) | 7.00 | 3.40 | Oblanceolate to spathulate | Obtuse | Tapering towards base | Entire | Distinct, thinning out towards apex | 4.00 | Emerge at 30-40°, dichotomizing and anastomosing to form meshes | Broad, smaller near midrib and narrow, elongated | 12-14 | 16–20 |
| Glossopteris retifera | Late Permian (Bijori/ Kamthi/ Raniganj formations) | 14.0 | 4.50 | Narrow oblanceolate to linear | Acute to obtuse | Acute to acute- cuneate | Entire | Persistent medium- broad, flat, evanescent towards apex | 3.00 | Emerge at 20–30°, run straight to margin | Broad, short, pentagonal- hexagonal | 6-9 | 12–18 |
| | Early Triassic (Panchet Formation) | 6.00 | 3.20 | Obovate | Acute | Indiscer- nible | Entire | Distinct, longitudinally ribbed | 2.00 | Emerge at 45°, dichotomizing and anastomosing | Oblong- polygonal meshes | 10-11 | 12–14 |

CURRENT SCIENCE, VOL. 107, NO. 10, 25 NOVEMBER 2014

namely size, shape, apex, base, midrib, angle of emergence of secondary veins, mesh type and average concentration of veins, the present analysis has been done on seven species of Glossopteris, viz. G. angustifolia, G. browniana, G. communis, G. conspicua, G. damudica, G. indica and G. retifera. These species continued their existence in Indian Gondwana sediments right from the Permian up to the early Triassic. Xeromorphic features, viz. thick cuticle, sunken stomata and stomatal papillae are usually found in the early Triassic plant taxa, e.g. Lepidopteris and Dicroidium. As evident from the fossil record, xeromorphic features have not been reported in any species of Glossopteris known from the early Triassic of India. However, xeromorphic features of plants cannot be interpreted purely in terms of possible antitranspirant properties. Certain features such as sunken stomata and stomatal papillae earlier regarded as antitranspirant in function, are now considered as adaptations for the growth of plants in high precipitation environments, serving a water-repellent role. Various xeromorphic features possibly served different functions and these properties act in combination, such as water repellence and resistance to toxic atmospheric gases and volcanic dust, etc. and may account for the development of xeromorphic features in a number of Mesozoic plants. Therefore, the use of these xeromorphic characters as indicators of palaeo-environmental aridity in the Mesozoic plant taxa needs supporting evidence from sedimentology, geochemical analysis, climate modelling and examination of modern ecological analogues⁴⁶.

The studied impression specimens of *Glossopteris* are from well-known late Permian and early Triassic



Figure 3. Size variation of species of *Glossopteris* from the late Permian and early Triassic sequence of India. *a*, *G. communis* from the late Permian (Raniganj Formation) of Raniganj Coalfield, West Bengal; BSIP specimen number II/629. *b*, *G. communis* from the early Triassic (Panchet Formation) of Deobar, Jharkhand; BSIP specimen number 35181/1246. *c*, *G. angustifolia* from the early Triassic (Panchet Formation) of Deobar, Jharkhand; BSIP Specimen Number 35180/1246 (scale bar = 1.0 cm).

CURRENT SCIENCE, VOL. 107, NO. 10, 25 NOVEMBER 2014

outcrops of peninsular Indian Gondwana basins, which are palynologically well dated. In the Gondwana sequence of Peninsular India, the Permian–Triassic Boundary is yet to be demarcated by absolute age



Figure 4. Graphical representation of variation of (*a*) laminar length, (*b*) laminar width and (*c*) midrib thickness in species of *Glossopteris* from late Permian and early Triassic sequences of India.

determination⁴⁷. However, the transition from Permian to Triassic is well documented by both plant megafossils and microfossils. High-resolution palynostratigraphy from the localities from where the specimens have been studied is well established. The late Permian palynoassemblage is characterized by *Densipollenites magnicorpus* zone and the early Triassic palynoassemblage is represented by *Falcisporites–Klausipollenites* zone^{48–51}.

The necessary measurements of the impressions of *Glossopteris* leaves were done from the original material. Illustrations of the already published papers of the original authors also have been taken into consideration. If any character of the leaf is not clear enough, it has been



Figure 5. Dimensions of the studied *Glossopteris* leaf specimens (length and width) in a cluster diagram.



Figure 6. Temporal variation of mean temperature during Permian and Triassic periods⁵⁵.

left as indeterminable. The results of the analysis are presented in Table 3. Evidences from the study conducted on the seven species of *Glossopteris* leaves clearly support the fact that there is indeed a reduction in the dimension of the lamina in the post-Permian sequence of India (Figure 3). The species from the early Triassic sequences are characterized by leaves possessing reduced laminar area with smaller length (Figure 4*a*) and width (Figure 4*b*) and narrower thickness of midrib (Figure 4*c*) in comparison to those of the underlying Permian sequences. Dimensions of the studied *Glossopteris* leaf specimens (length and width) have been given in the form of a cluster diagram (Figure 5).

Modern-day high-resolution regional palaeoecological studies have proved the myth wrong, which stated that the mass extinction event had little macroecological or evolutionary consequence for terrestrial plants³⁸. The early Triassic (Panchet Formation) in India witnessed more arid or semi-arid climate in comparison to the Permian. The early Triassic experienced greenhouse conditions with a warmer phase (Figure 6) due to global rise of temperature coupled with episodes of intense volcanism⁵²⁻⁵⁶. The reduced size of the lamina is one of the strategies adapted by the plants during the adverse condition in early Triassic when there was indeed a shortage of essential nutrients in the soil, in addition to seasonal dry climate, irregular rainfall and widespread aridity. The leaves possessed cuticles with sunken stomata since the atmosphere during the early Triassic had higher levels of CO_2 and lower O_2 levels^{57–60} (Figure 7). The *Glossopteris* flora of late Permian was adapted to temperate, cool and moist environments^{60,61}. The phenomenon of dwarfism as evidenced in Glossopteris has been observed only up to the early Triassic, which is represented by the Panchet Formation. During the middle to late Triassic, typical Dicroidium flora associated with Lepidopteris takes over the preceding Glossopteris flora.

Records of plant fossils substantiate the evidence that numerous physiological, reproductive and behavioural traits enabled the smaller sized plant species to persist in extreme climatic conditions^{38,62–66}. With the onset of new



Figure 7. Temporal variation in atmospheric CO_2 and O_2 during the Permian and Triassic periods^{56,58}.

environmental conditions during the early Triassic, surpassing the major extinction event, the smaller sized features of glossopterids became prevalent and continued their existence, though, during the late Triassic they gradually became extinct.

- 1. Brongniart, A., Sur la classification et la distribution des végétaux fossiles on géneral, et sur ceux des terrains de sédiment supérieur en particulier. Mus. Hist. Nat. Paris, Mém., 1822, 8, 203-348.
- 2. Brongniart, A., Histoire des végétaux fossiles on researches botaniques et geologiques sur les végétaux renfermés dans les diverses couches du globe, Paris, 1828-37, pp. 1-488.
- 3. Arber, E. A. N., Catalogue of the Fossil Plants of the Glossopteris Flora, British Museum, London, 1905, p. 295.
- 4. Chandra, S. and Surange, K. R., Revision of the Indian Species of Glossopteris, Birbal Sahni Institute of Palaeobotany, Lucknow, India, Monograph 2, 1979, pp. 1–291.
- 5. Kovács-Endrödy, E., Notes on Glossopteris Angustifolia Brongniart. Palaeobotanist, 1981, 28-29, 53-62.
- 6. McLoughlin, S., Glossopteris insights into the architecture and relationships of an iconic Permian Gondwanan plant. J. Bot. Soc. Bengal, 2011, 65, 93-106.
- 7. Delevoryas, T. and Person, C. P., Mexiglossa varia gen. et sp. nov., a new genus of glossopteroid leaves from the Jurassic of Oaxaca, Mexico. Palaeontogr. A, 1975, 154, 114-120.
- 8. Srivastava, P. N., Studies in the Glossopteris Flora of India 4. Glossopteris, Gangamopteris and Palaeovittaria from the Raniganj Coalfield. Palaeobotanist, 1957, 5, 1-45.
- 9. Goswami, S., Singh, K. J. and Chandra, S., Palaeobotany of Gondwana basins of Orissa State, India: a bird's eye view. J. Asian Earth Sci., 2006, 28, 218-233.
- 10. Goswami, S., Singh, K. J. and Chandra, S., Pteridophytes from Lower Gondwana formations of the Ib River Coalfield, Orissa and their diversity and distribution in the Permian of India. J. Asian Earth Sci., 2006, 28, 234-250.
- 11. Goswami, S., Record of Lower Gondwana megafloral assemblage from Lower Kamthi Formation of Ib River Coalfield, Orissa, India. J. Biosci., 2006, 31, 115-128.
- 12. Singh, K. J., Goswami, S. and Chandra, S., First report of genus Gangamopteris from Gondwana sediments of Ib-River Coalfield, Orissa. J. Geol. Soc. India, 2006, 68, 893-905.
- 13. Singh, K. J., Goswami, S. and Chandra, S., The genus Glossopteris from Lower Gondwana Formations of Ib-River Coalfield, Orissa, India. J. Palaentol. Soc. India, 2006, 51, 81-107.
- 14. Pal, P. K., Srivastava, A. K. and Ghosh, A. K., Plant fossils of Maitur Formation: possibly the ultimate stage of Glossopteris flora in Raniganj Coalfield, India. Palaeobotanist, 2010, 59, 33-45.
- 15. Srivastava, A. K. and Agnihotri, D., Upper Permian plant fossil assemblage of Bijori Formation: a case study of Glossopteris flora beyond the limit of Raniganj Formation. J. Geol. Soc. India, 2010, 76 47-62
- 16. Cúneo, R., Isbell, J., Taylor, E. L. and Taylor, T. N., The Glossopteris flora from Antarctica: taphonomy and paleoecology. Comptes Rendus XII ICC-P, Buenos Aires, 1993, 2, 13-40.
- 17. McLoughlin, S., Lindstorm, S. and Drinnan, A. N., Gondwanan floristic and sedimentological trends during the Permian-Triassic transition: new evidence from the Amery Group, northern Prince Charles Mountains, East Antarctica. Antarctic Sci., 1997, 9, 281-298
- 18. Bose, M. N., Triassic floras. In Aspects and Appraisals of Indian Palaeobotany (eds Surange, K. R. et al.), Birbal Sahni Institute of Palaeobotany, Lucknow, 1974, pp. 285-293.
- 19. Banerji, J., Maheshwari, H. K. and Bose, M. N., Some plant fossils from the Gopad River Section near Nidpur, Sidhi District, Madhya Pradesh. Palaeobotanist, 1974, 23, 59-71.

RESEARCH COMMUNICATIONS

- 20. Bose, M. N. and Banerji, J., Some fragmentary plant remains from the lower Triassic of Auranga Valley, District Palamau. Palaeobotanist, 1974, 23, 139-146.
- 21. Banerji, J. and Bose, M. N., Some lower Triassic plant remains from Assansol region, India. Palaeobotanist, 1975, 24, 202-212.
- 22. Urbanek, A., Biotic crises in the history of Upper Silurian graptoloids: a paleobiological model. Hist. Biol., 1993, 7, 29-50.
- 23. Harris, P. J., Kauffman, E. G. and Hansen, T. A., Models for biotic survival following mass extinctions. In Biotic Recovery from Mass Extinction Events (ed. Hart, M. B.), Geological Society, London, 1996, vol. 102, pp. 41-60.
- 24. Harris, P. J. and Knorr, P. O., What does the 'Lilliput effect' mean. Palaeogeogr., Palaeoclimatol., Palaeoecol., 2009, 284, 4-10.
- 25. Chen, Z. Q. and Benton, M. J., The timing and pattern of biotic recovery following the end-Permian mass extinction. Nature Geosci 2012 5 375-383
- 26. Stanley, S. M., An explanation for Cope's rule. Evolution, 1973, 27, 1-26.
- 27. Donovan, S. K., Mass Extinctions: Processes and Evidence, Columbia University Press, New York, 1989, pp. 1-266.
- 28. Witting, L., Dwarfing and Extinction. A General Theory of Evolution, Peregrine Publisher, Denmark, 1997, pp. 1-348.
- 29. Wade, B. S. and Twichett, R., Extinction, dwarfing and the Lilliput Effect. Palaeogeogr., Palaeoclimatol., Palaeoecol., 2009, 284, 1-3.
- 30. Payne, J. L., Evolutionary dynamics of gastropod size across the end-Permian extinction and through the Triassic recovery interval. Palaeobiology, 2005, 31, 269-290.
- 31. Twitchett, R. J., Incompleteness of the Permian-Triassic fossil record: a consequence of productivity decline? Geol. J., 2001, 36, 341-353
- 32. Mutter, R. J. and Neuman, A. G., Recovery from the end-Permian extinction event: evidence from 'Lilliput Listracanthus'. Palaeogeogr., Palaeoclimatol., Palaeoecol., 2009. 284, 22-28.
- 33. Metcalfe, B., Twitchet, R. and Price-Lloyd, N., Size and growth rate of 'Lilliput' animals in the earliest Triassic. Palaeogeogr., Palaeoclimatol., Palaeoecol., 2011, 308, 171-180.
- 34. Song, H., Tong, J. and Chen, Z. Q., Evolutionary dynamics of the Permian-Triassic foraminifer size: evidence for Lilliput effect in the end-Permian mass extinction and its aftermath. Palaeogeogr., Palaeoclimatol., Palaeoecol., 2011, 308, 98-110.
- 35. Erwin, D. H., The Great Paleozoic Crisis: Life and Death in the Permian, Columbia University Press, New York, 1993, p. 327.
- 36. Lockwood, R., Body size, extinction events, and the early Cenozoic record of veneroid bivalves: a new role for recoveries? Palaeobiology, 2005, 31, 578-590.
- 37. Wade, B. S. and Pearson, P. N., Planktonic foraminiferal turnover, diversity fluctuations and geochemical signals across the Eocene/Oligocene boundary in Tanzania. Mar. Micropaleontol., 2008.68.244-255.
- 38. McElwain, J. C. and Punyasena, S. W., Mass extinction events and the plant fossil record. Trends Ecol. Evol., 2007, 22, 548-557.
- 39. McElwain, J. C., Beerling, D. J. and Woodward, F. I., Fossil plants and global warming at the Triassic-Jurassic boundary. Science, 1999, 285, 1386-1390.
- 40. Bacon, K. L., Belcher, C. M., Haworth, M. and McElwain, J. C., Increased atmospheric SO2 detected from changes in leaf physiognomy across the Triassic-Jurassic boundary interval of East Greenland. PLoS One, 2013, 8, e60614.
- 41. Willis, K. J. and McElwain, J. C., The Evolution of Plants, Oxford University Press, UK, 2001, pp. 239-243.
- 42. Looy, C. V., Brugman, W. A., Dilcher, D. L. and Visscher, H., The delayed resurgence of equatorial forests after the Permian-Triassic ecological crisis. Proc. Natl. Acad. Sci. USA, 1999, 96, 13857-13862
- Looy, C. V., Twitchett, R. J., Dilcher, D. L., Van Konijnenburg-43. Van Cittert, J. H. A. and Visscher, H., Life in the end-Permian dead zone. Proc. Natl. Acad. Sci. USA, 2001, 98, 7879-7883.

- Benton, M. J. and Newell, A. J., Impacts of global warming on Permo-Triassic terrestrial ecosystems. *Gondwana Res.*, 2012; doi: 10.1016/j.gr.2012.12.010.
- Sastry, M. V. A. *et al.*, Stratigraphic lexicon of Gondwana Formations of India. *Geol. Surv. India*, *Miscl. Publ.*, 1977, 36, 1–170.
- Haworth, M. and McElwain, J. C., Hot, dry, wet, cold or toxic? Revisiting the significance of leaf and cuticular micromorphology. *Palaeogeogr.*, *Palaeoclimatol.*, *Palaeoecol.*, 2008, 262, 79–90.
- 47. Sarkar, A., Yoshioka, H., Ebihara, M. and Naraoka, H., Geochemical and organic carbon isotope studies across the continental Permo-Triassic boundary of Raniganj basin, Eastern India. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 2003, **191**, 1–14.
- Vijaya and Tiwari, R. S., Role of spore-pollen species in demarcating the Permo-Triassic boundary in Raniganj Coalfield, West Bengal. *Palaeobotanist*, 1986, **35**, 242–248.
- Tiwari, R. S. and Tripathi, A., Palynological zones and their climate inference in the coal bearing Gondwana of Peninsular India. *Palaeobotanist*, 1988, 36, 87–101.
- 50. Srivastava, S. C. and Jha, N., Palynostratigraphy and correlation of Permian–Triassic sediments in Budharam Area, Godavari Graben, India. J. Geol. Soc. India, 1995, **46**, 647–653.
- Srivastava, S. C., Anand-Prakash and Kar, R., Palynology of Permian–Triassic sequence in Iria Nala, Tatapani–Ramkola Coalfield, India. *Palaeobotanist*, 1997, 46, 75–80.
- Srivastava, S. C. and Prakash, N., Climatic fluctuation based on two hundred million year old (± 225 M.Y.-Triassic) plants in Indian peninsula: a significant advance in botany. In *Glimpses in Botany* (ed. Mukherji, K. G.), A.P.H. Publishing Corporation, New Delhi, 2000, pp. 80–95.
- 53. Frakes, L. A., Francis, J. E. and Syktus, J. I., *Climate Models of the Phanerozoic*, Cambridge University Press, Cambridge, 1992, p. 274.
- 54. Retallack, G. J. and Zarza, A., Middle Triassic paleosols and paleoclimatology of Antarctica. J. Sediment. Res., 1998, **68**, 169–184.
- Scotese, C. R., Boucot, A. J. and McKerrow, W. S., Gondwanan palaeogeography and palaeoclimatology. J. Afr. Earth Sci., 1999, 28, 99–114.
- Berner, R. A., Examination of hypotheses for the Permo-Triassic boundary extinction by carbon cycle modelling. *Proc. Natl. Acad. Sci. USA*, 2002, 99, 4172–4177.
- 57. Ward, P., Impact from the deep. Sci. Am., 2006, 295(4), 64-71.
- Berner, R. A., GEOCARBSULF: a combined model for Phanerozoic atmospheric O₂ and CO₂. *Geochim. Cosmochim. Acta*, 2006, 70, 5653–5664.

- Sues, H.-D. and Fraser, N. C., *Triassic Life on Land: The Great Transition*, Columbia University Press, New York, 2010, p. 179.
- McLoughlin, S., Glossopterid megafossils in Permian Gondwanic non-marine biostratigraphy. In *Gondwana Eight: Assembly, Evolution and Dispersal* (eds Findlay, R. H. *et al.*), Balkema, Rotterdam, 1993, pp. 253–264.
- McManus, H. A., Taylor, E. A., Taylor, T. N. and Collinson, J. W., A petrified *Glossopteris* flora from Collinson Ridge, central Transarctic Mountains: Late Permian or Early Triassic? *Rev. Palaeobot. Palynol.*, 2002, **120**, 233–246.
- Knoll, A. H., Patterns of extinction in the fossil record of vascular plants. In *Extinctions* (ed. Nitecki, M.), University of Chicago Press, Chicago, 1984, pp. 21–68.
- Rees, P. M., Land-plant diversity and the end-Permian mass extinction. *Geology*, 2002, 30, 827–830.
- Wing, S. L., Mass extinctions in plant evolution. In *Extinctions in the History of Life* (ed. Taylor, P. T.), Cambridge University Press, Cambridge. 2004, pp. 61–97.
- 65. Kerp, H. A., Hamad, A., Vörding, B. and Bandel, K., Typical Triassic Gondwanan floral elements in the Upper Permian of the paleotropics. *Geology*, 2006, **34**, 265–268.
- Xiong, C. and Wang, Q., Permian–Triassic land-plant diversity in South China: was there a mass extinction at the Permian/Triassic boundary? *Palaeobiology*, 2011, 37, 157–167.
- Roy Chowdhury, H. K., Sastry, M. V. A., Shah, S. C., Singh, G. and Ghosh, S. C., Triassic floras in India. In *Gondwana Geology* (ed. Campbell, K. S. W.), Third-International Gondwana Symposium, Australian National University Press, Canberra, 1973, pp. 149–159.
- Datta, P., Gondwana lithostratigraphy of Peninsular India. Gondwana Res., 2002, 5, 540–553.

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